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Validation of the HITRAN 2016 and GEISA 2015 line lists using ACE-FTS solar occultation observations



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ABSTRACT

The ExoMars Trace Gas Orbiter (TGO) began its nominal science phase at Mars in April 2018, following releases of editions to two major spectroscopic line lists: GEISA 2015 (Gestion et Etude des Informations Spectroscopiques Atmosphériques: Management and Study of Atmospheric Spectroscopic Information), and HITRAN 2016 (High Resolution Transmission). This work evaluates both line lists over the spectral region between 2325–4350 cm⁻¹ using terrestrial solar occultation observations made by the Atmospheric Chemistry Experiment Fourier Transform Spectrometer (ACE-FTS). This spectral region is targeted on Mars by two complementary solar occultation instruments on TGO that will monitor temperature and pressure, aerosols, the abundance of CO_2 , CO_1 , H_2O_1 , HDO, CH_4 , and other undetected trace gases. Major updates to GEISA 2015 and HITRAN 2016, with respect to previous editions, have been focused on CO₂ absorption features in support of Earth-observing missions to monitor greenhouse. Since CO₂ is the dominant absorber on Mars, making up 96.5% of the atmosphere, validating the updated line lists is critically important before their deployment for ExoMars. We report that updated CO₂ parameters make significant improvements to spectral fits made when using both line lists. Several updates to H₂O lines in both line lists also show improvement. The primary difference we observe between the two line lists comes from O₃ absorption features near 3850 cm⁻¹ and from several CH₄ absorption lines in the regions 2800-3200 cm⁻¹ and 4000-4300 cm⁻¹. Because of these differences, we find that using HITRAN 2016 tends to result in better spectral fits, especially below 30 km, than using GEISA 2015 in this spectral region. Differences are strongly reduced with increasing altitude (>40 km) as pressure and gas abundance falls off. It was also discovered that several new errors in both new editions of GEISA and HITRAN were introduced since the HITRAN 2012.

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1. Introduction

New editions of two major spectroscopic line lists have been recently released: Gestion et Etude des Informations Spectroscopiques Atmosphériques: Management and Study of Atmospheric Spectroscopic Information (GEISA) in 2015 [30]; and High Resolution Transmission (HITRAN) in 2016 [26]. Here, we present a comparison of spectral fits to solar occultation measurements of the

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Earth's atmosphere made by the Canadian Space Agency's (CSA's) Atmospheric Chemistry Experiment (ACE) Fourier transform spectrometer (FTS) using both line lists. This work was motivated by the arrival of the European Space Agency (ESA) and Roscosmos' ExoMars Trace Gas Orbiter (TGO) at Mars in October 2016. The TGO carries two infrared remote sensing instrument suites, the Atmospheric Chemistry Suite (ACS) [33] and the Nadir and Occultation for Mars Discovery (NOMAD) [65]. Both instrument suites carry channels dedicated to making solar occultation observations at Mars in the infrared wavenumber range of 2325–4350 cm⁻¹.

The controversial observation of methane (CH₄) in the Martian atmosphere [23,25,36,46,68,69] is one of the key motivations of the ExoMars mission [33,66,74]. NOMAD and ACS will search for CH₄ by making solar occultation observations of its v_2

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vibration-rotation band centred near 3000 cm⁻¹, which hosts the strongest absorption features available to both instruments. A key benefit of searching for CH₄ in this region is that it is relatively clear of interfering absorption lines from CO₂. There is a recently-observed, weak vibration-rotation band of carbon dioxide (CO₂) that overlaps CH₄ in this region [9,67] and one of our specific objectives was to evaluate its spectroscopic parameters in GEISA 2015 and HITRAN 2016.

ACS and NOMAD will also make detailed measurements of water vapour (H_2O), having the capability of distinguishing isotopologues. The ratio of HDO to H_2O has been used as a critical indicator of the Martian climate in the ancient past [e.g., 22, 35]. The relative abundance of HDO on Mars is enriched relative to Earth, which supports mechanisms for hydrogen escape from the atmosphere, which are preferential towards the lighter isotope [14]. ACS and NOMAD will be able measure the vertical structure of H_2O isotopologues, and monitor their ratios seasonally, spatially, and vertically.

In support of the study of Earth's contemporary climate, and the carbon cycle, with the Orbiting Carbon Observatory (OCO-2) [18,19], the Greenhouse gases Observing SATellite (GOSAT) [72], and the Total Carbon Column Observing Network (TCCON) [71], a large effort has been undertaken to refine and reduce the uncertainty of the spectroscopic parameters of CO_2 . While the surface pressure of Mars is only between 550–720 Pa, roughly 0.5–0.7% of Earth's, the volume mixing ratio of CO_2 is 0.965, or 2400 times that on Earth (400 ppmv). Therefore, CO_2 absorption lines in Mars solar occultation spectra will be deeper and broader than for spectra recorded at Earth, and small changes in the spectroscopic parameters may have a large impact on trace gas retrievals made at Mars.

Our goal was to evaluate the new CO₂, H₂O, and CH₄ parameters by looking at whether their impact when fitting terrestrial spectra was positive or negative. We have performed spectral fitting for 125 sets of ACE-FTS occultation spectra (resolution of 0.02 cm⁻¹) over 50 spectral windows covering the spectral range of 2430–4450 cm⁻¹ using the HITRAN 2012, HITRAN 2016, and GEISA 2015 line lists.

HITRAN 2016 and GEISA 2015 are compilations of data sources, many of which are shared. One of our most important results is that the updated CO_2 parameters for some of the stronger vibration-rotation bands, especially those centred at 3550 cm⁻¹ result in strongly improved spectral fits when using either HITRAN 2016 or GEISA 2015. However, we have also found that there are large differences in the spectroscopic parameters of ozone (O_3), CH₄, and H₂O between the two data sets, and that using HITRAN 2016 leads to improved spectral fits compared to GEISA 2015. This result is significant since methane is one of the strongest absorbers in the Earth's atmosphere, and is one of the most variable gases. Methane is of key importance for TGO, as ACS and NOMAD both aspire to make its irregular and controversial detection in the Martian atmosphere definitive. However, at this time no methane features have been observed at Mars by TGO instruments [34].

While individual contributions to the line lists are validated, they are often done in laboratory settings, observing controlled gas samples, rather than with observations of an atmosphere [e.g., 26, 30, and references therein]. [3] previously made a direct comparison between H₂O transitions in older versions of GEISA and HI-TRAN, and qualitatively showed their differences using modelled spectra for Venus at high temperatures and above 4000 cm⁻¹. A comprehensive validation of the GEISA line list was done using TC-CON and the Infrared Atmospheric Sounding Interferometer (IASI) [15] by Armante et al. [1]. They describe a technique used to determine whether spectroscopic parameters should be used to update the GEISA database based on comparisons of computed spectra to observations. They specifically show H₂O and HDO in the ExoMars

region of interest and highlight improvements since GEISA 2011. They also compare CH_4 lines above 6000 cm⁻¹ to HITRAN 2012 and note an improvement to the residuals. This method was used in the compilation of GEISA 2015 [30].

A comprehensive validation of the HITRAN 2012 line list was undertaken by Toon et al. [60] using the MkIV balloon-borne FTS [59]. They divided the spectral region between 670–5620 cm⁻¹ into fitting windows and quantitatively evaluated the best-fit residuals across the spectral range and with altitude for several versions of HITRAN released since 2000. They noted specific errors in the data base, where improvements were made, and where previous versions performed better. Their work was influential on the compilation of the latest version of HITRAN evaluated here [26]. Updates to [60] are included in [58] and include evaluation of HITRAN 2016 and the TCCON internal line list, validation with laboratory spectra, and a specific analysis of CO_2 features. This work follows [60] by using a similar quantitative evaluation technique and covering part of the same spectral region.

2. HITRAN 2016

The HITRAN (high-resolution transmission molecular absorption) database was first compiled for the Air Force Geophysics Laboratory (AFGL) by McClatchey et al. [44] and major updated editions have been released on a four year cycle since 1992 [51]. The 2016 version of the HITRAN database [26] describes changes made since the 2012 edition [52]. Among the most significant additions to the database have been the inclusion of spectroscopic parameters for collision-induced broadening from non-nitrogen based atmospheres and for non-Voigt line profiles [70]. The need for line broadening parameters in atmospheres primarily composed of gases other than N₂ has been motivated by extra-terrestrial spectroscopic applications, e.g., Mars, which is 96% CO₂. HITRAN 2016 includes a sparse set of broadening parameters for atmospheres composed of H₂, He, or CO₂ for a subset of gases that includes CO, OCS, SO₂, NH₃, HF and HCl, which are all sought at Mars by the ExoMars TGO. For very high-resolution applications, HITRAN 2016 also includes parameters for the speed-dependent Voigt, Galatry, and Hartman-Tran line shapes. These are again only available for a subset of wavenumbers and only for H₂O, CO, H₂, O₂, N₂O, HF and HCl. Because of the complexity of the newly included parameters, newly-developed online tools are now used to create userdefined database versions. The 2016 version also expands the list of available molecules with the addition of C₂N₂ and COCl₂. While the number of additional lines, and the number of lines for which more accurate measurements have been made, is vast, we will primarily focus on key species relevant to Earth and Mars: CO₂, H₂O and CH₄.

For CO₂, [26] identifies the imperative for high-accuracy spectroscopic parameters driven by GOSAT, OCO-2, TCCON, and others, and identifies validated improvements between the 2008 and 2012 versions of the HITRAN database [60]. The 2012 version of the database was largely built on theoretical fits of the effective Hamiltonian or effective dipole moments, compiled as the CDSD-296 database [56], supplemented by higher-accuracy experimental measurements made by Toth et al. [61-63]. The line intensity calculations in CDSD-296 have high uncertainties ($\sim 20\%$) and two new sets of theoretical computation have been produced: the Ames list [27] and the UCL-IAO list [73]. These have been extensively validated experimentally, and [26] refers to 14 such studies that show that the UCL-IAO list tends to be more accurate, and that the uncertainties for the 2016 version of HITRAN can be pushed down to the order of 0.5%. However, the majority of the experimental work focuses on important CO₂ bands for Earth observation (e.g., for OCO-2) that lie outside the range of high-resolution solar occultation experiments on TGO, above 5000 cm⁻¹ (near 1.6

and 2 µm). Only laboratory measurements presented in [43] and [21] cover the 2300–4400 cm⁻¹ range, and the latter only does so near 3730 cm⁻¹. The 2016 HITRAN line list for CO₂ between 2300–4400 cm⁻¹ is a combination of CDSD-296 theoretical calculations [56], UCL-IAO or Ames theoretical calculations where better or newly available [27,73], and laboratory measurements where available and with low enough uncertainty [e.g., 21, 43, 63].

For H₂O, the HITRAN 2012 line list was made up of *ab initio* calculations that comprised the BT2 line list [7], with updates, where available, from calculations using a more accurate method [39,40]. Newer calculations using the methodology of [39] have been made as part of an effort by the International Union of Pure and Applied Chemistry (IUPAC) task group [57, and references therein]. The new calculations have been validated experimentally by Birk et al. [10] and the results have been used to update both the IU-PAC database and the HITRAN line list. Extensive laboratory measurements have also been made for the German Aerospace Agency (DLR) in the spectral range between 1850–4000 cm⁻¹ by Loos et al. [41,42]. When available, these replace the calculated line strengths of IUPAC and [39].

The CH₄ data in HITRAN 2012 was comprised of the data set described in [12], which was a combination of theoretical calculations and experimental measurements. This data set replaced over 70% of the HITRAN 2008 line list for CH₄ [53]. However, analysis of high-resolution solar occultation measurements in the Earth's atmosphere made by the MkIV interferometer [59] determined that there were still several errors and omissions in the HITRAN 2012 CH₄ data [60], especially in the spectral region of the ν_2 transition critical to ExoMars, near 3000 cm⁻¹. Errors that were identified were replaced by either the HITRAN 2008 values, or computations made by [64]. Several laboratory studies have been recently undertaken, but the results have not yet been incorporated into the HITRAN line list, but an update to the 2016 edition is expected in the interim [26].

3. GEISA 2015

The GEISA line list was first compiled in the early 1970s at the Laboratoire de Météorologie Dynamique (LMD) to support their radiative transfer investigations [13]. Key motivations for the compilation were to include new gases important for planetary atmospheric applications, and to co-develop software tools to easily use the database. One distinguishing feature is to treat certain isotopologues with distinct symmetries as independent species (such as HDO for H₂O and CH₃D for CH₄) [30]. Comparing the available gases in current versions of GEISA and HITRAN, GeH₄, C₃H₈, C₃H₄, and C₆H₆ are unique to GEISA, while HOBr, O, H₂, and CS are unique to HITRAN. There are also several minor isotopologues of trace gases unique to both. Updates to GEISA are made after evaluating the relevance of new data, the efficiency of including it, and after undergoing a validation process as described in [1].

For CO₂, the GEISA 2011 database was replaced by the CDSD-296 database [56]. CDSD-296 is also the primary source of CO₂ parameters in the 2016 edition of HITRAN. GEISA also contains three isotopologues not contained in CDSD-296 from laboratory measurements by Jacquemart et al. [29,43] (and others at higher wavenumbers than 4400 cm⁻¹).

Extensive updates to H_2O were made empirically for GEISA 2015 by a consortium of eight laboratories, nearly tripling the number of available lines since the 2011 edition. In the spectral region of interest to ExoMars ($\sim 2300-4400 \text{ cm}^{-1}$), these measurements were made by the Laboratorie Inter-Universitaire des Systèmes Atmosphériques (LISA), the Institute of Atmospheric Optics (IAO), and University College, London [30]. Updates to $H_2^{16}O$ come from Coudert et al. [17], and updates for $H_2^{17}O$ and $H_2^{18}O$ come from

Lodi et al. [39,40] and the IUPAC efforts, which is the same source as for HITRAN 2016. These were supplemented or updated by measurements made by [16]. GEISA 2015 also newly includes lines for two isotopologues of D₂O not included in HITRAN.

Updates to CH_4 in this spectral range mainly come from the work of [47,48] which use the same methodology as [12] (HITRAN 2012). The validation work of [1] showed some imprecision in the new parameters, resulting in some CH_4 lines from GEISA 2011 being retained.

4. Methodology

In this study, we break the wavenumber range of the ExoMars solar occultation spectrometers up into discrete fitting windows and analyze terrestrial solar occultation spectra recorded by ACE-FTS using the Jet Propulsion Laboratory Gas Fitting (GFIT or GGG) software suite. 125 occultations were analyzed. During an occultation, a series of observations of the sun are made while the limb of the atmosphere lies between the solar disk and the instrument. For each window, residuals were computed for each altitude level. The means of the residuals were taken at levels of equal pressure, and the root-mean-square (RMS) and standard deviation (σ) were computed for each fitting window. The means of the residuals were taken, rather than computing the residuals of mean spectra, due to variations in line depths between occultations, especially for CH₄ lines. In general, the lower the results RMS of the mean residuals is, the more accurate the spectroscopic parameters used in the fitting are. This methodology is very similar to that used by Toon et al. [60] who analyzed MkIV spectra.

ACE-FTS is a compact, double-pass interferometer with a spectral resolution of 0.02 cm⁻¹ and a spectral range of 750–4400 cm⁻¹. It has been operating continuously in low-Earth orbit since 2003. The 125 ACE-FTS occultations analyzed were recorded between 2004 and 2012 and are unrestricted in longitude and season. Most observations are at high latitudes due to the ACE orbit (650 km with an inclination of 74°). A sequence of measurements is made with an observation every 1–6.5 km, and on average every 4 km. Two detectors provide an simultaneous spectral range of 750–4400 cm⁻¹ [8].

Spectral fitting in this study was done with GGG, which is also being used to analyze solar occultation observations made by ACS on ExoMars. GGG is developed from early Occultation Display Spectra (ODS) used by the ATMOS FTS that flew on the space shuttle [28,49]. It is a robust software suite adaptable for solar occultations [50,59,60], ground based observations [71], or laboratory measurements. For each altitude and each fitting window, GGG computes a spectrum from a set of parameters that include the calculated optical path, vertical profiles of pressure and temperature, and a priori gas volume mixing ratio (VMR) vertical profiles. GGG then performs non-linear least squares fitting to adjust VMR scaling factors (as well as other optional parameters such as continuum level and frequency shift) to obtain a best fit. Retrieved VMR vertical profiles can be obtained by inverting the matrix of VMR scale factors for each target gas at each altitude, with the matrix of slant paths [54,71].

A priori temperature, pressure, and specific humidity vertical profiles were derived from National Centers for Environmental Prediction (NCEP) reanalysis data [32] up to 40 km. The US standard atmosphere was used above 40 km.

For this work, we consider the fitting residuals for each window and focus on three pressure levels where we take the means of the residuals: 0.052, 0.0029, and 0.00023 atm, corresponding to 20, 40, and 60 km respectively. For each occultation, fitting was performed at all altitudes and in all spectral windows. The closest altitude level to the predetermined pressure levels was identified.

Table 1

List of spectral fitting windows used to evaluate the HITRAN 2016 and GEISA 2015 line lists with ACE-FTS solar occultation observations of the Earth's atmosphere. Given for each fitting window are: centre wavenumber, window width, and gases fit in the window.

Centre $\tilde{\nu}$ (cm ⁻¹)	Width (cm ⁻¹)	Gases fit
2455.5	25.9	CO ₂ , CH ₄ , H ₂ O, N ₂ O, O ₃
2491.5	47.9	CO ₂ , CH ₄ , H ₂ O, N ₂ O, O ₃
2551.55	71.35	H ₂ O, CO ₂ , O ₃ , N ₂ O, CH ₄
2615.74	56.2	CO ₂ , CH ₄ , H ₂ O, N ₂ O, O ₃ , HCl, C ₂ H ₆
2666.6	44.76	H ₂ O, CO ₂ , O ₃ , N ₂ O, CH ₄ , HCl, C ₂ H ₆
2689.8	0.62	HDO, CO ₂ , O ₃ , N ₂ O, CH ₄
2692.76	0.55	HDO, CO ₂ , O ₃ , N ₂ O, CH ₄
2708.17	0.54	HDO, CO ₂ , O ₃ , N ₂ O, CH ₄
2722.25	52.0	CO ₂ , CH ₄ , C ₂ H ₆ , H ₂ O, N ₂ O, O ₃ , HCl, HDO
2780.74	65.15	CO ₂ , CH ₄ , C ₂ H ₆ , H ₂ O, N ₂ O, O ₃ , HCl, HDO
2801.6	0.49	HDO, CO ₂ , O ₃ , N ₂ O, CH ₄
2825.0	22.95	H ₂ O, O ₃ , N ₂ O, CH ₄ , HDO, HCl
2849.15	24.55	HDO, CO ₂ , H ₂ O, CH ₄ , C ₂ H ₆ , O ₃ , N ₂ O, HCl
2904.43	85.10	H ₂ O, O ₃ , N ₂ O, CH ₄ , HDO, HCl, NO ₂ , OCS
2973.65	1.21	C ₂ H ₆ , CO ₂ , H ₂ O, CH ₄ , O ₃
2978.2	62.0	CO ₂ , CH ₄ , C ₂ H ₆ , C ₂ H ₄ , H ₂ O, O ₃ , HCl, HDO
2983.49	0.78	C ₂ H ₆ , CO ₂ , O ₃ , CH ₄
2986.74	0.58	C_2H_6 , CO_2 , O_3
2989.98	0.80	C ₂ H ₆ , CO ₂ , O ₃ , CH ₄
2993.52	1.23	C ₂ H ₆ , CO ₂ , O ₃ , CH ₄ , H ₂ O
3022.13	18.14	CH ₄ , H ₂ O, C ₂ H ₆ , C ₂ H ₄ , O ₃ , HCl
3035.08	9.2	CH_4 , H_2O , C_2H_6 , C_2H_4 , O_3
3057.72	5.1	CH ₄ , H ₂ O, O ₃
3065.86	3.35	CH ₄ , H ₂ O, O ₃
3077.36	2.54	CH_4 , H_2O , O_3

The residuals were computed for each window and the RMS and standard deviation were computed from the residual as:

$$RMS^{2} = \sum x_{i}^{2}/N,$$

$$\sigma^{2} = \sum (x_{i} - \bar{x})^{2}/(N - 1),$$
(1)

where *N* is the number of spectral points, x_i is the residual value at the *i*th spectral point, and \bar{x} is the mean of the residual. In general, for this application $\bar{x} \sim 0$ and RMS² $\sim \sigma^2$. For strong deviations between the computed and observed spectra, σ^2 will reflect the magnitude of deviations from the mean of the residuals, while the RMS will reflect the deviation from zero, which is the expected outcome for a good fit. Therefore, the presented results will use the RMS values.

The fitting windows used are given in Tables 1 and 2, and panel a in Fig. 1 shows their distribution over a sample of ACE-FTS solar occultation transmission spectrum. Panels b-e in Fig. 1 illustrate the locations and magnitudes of absorption features due to major gas species in the Earth's atmosphere: CO_2 , H_2O , CH_4 , O_3 , and N_2O .

5. Results

We begin with an overview of the differences between the HI-TRAN 2016 and GEISA 2015 line lists in the ExoMars spectral region of interest, 2325–4350 cm⁻¹. Fig. 2 shows mean ACE-FTS spectra averaged at three different pressure levels corresponding to approximately 20, 40, and 60 km tangent altitudes. Each primary panel, for each pressure level, shows the mean ACE-FTS spectra, averaged over 125 occultations, for each fitting window (Fig. 2 does not show continuous ACE-FTS spectra). Shown over top of the mean ACE-FTS spectra, is the mean of the calculated best fits for each window when using the HITRAN 2016 line list (as an example). It is important to note that the noise level of the ACE-FTS spectra is not constant with wavenumber [11], and the noise can

Table 2

List of spectral fitting windows used to evaluate the HITRAN 2016 and GEISA 2015 line lists with ACE-FTS solar occultation observations of the Earth's atmosphere. Given for each fitting window are: centre wavenumber, window width, and gases fit in the window.

Centre $\tilde{\nu}$ (cm ⁻¹)	Width (cm ⁻¹)	Gases fit
3089.75	17.5	CH ₄ , H ₂ O, C ₂ H ₆ , C ₂ H ₄ , CO ₂ , O ₃
3126.65	58.0	CO ₂ , CH ₄ , C ₂ H ₄ , H ₂ O, O ₃ , HDO
3202.0	90.0	H ₂ O, CO ₂ , CH ₄ , O ₃ , HCN
3292.0	90.0	H ₂ O, CO ₂ , CH ₄ , O ₃ , HCN, N ₂ O
3391.15	108.6	CO ₂ , H ₂ O, C ₂ H ₂ , NH ₃ , HCN, N ₂ O, O ₃ , HDO
3489.0	86.0	H ₂ O, CO ₂ , O ₃ , N ₂ O
3577.0	86.0	H ₂ O, CO ₂ , O ₃ , N ₂ O, HNO ₃ , HDO
3665.0	86.0	H ₂ O, CO ₂ , O ₃ , N ₂ O, HNO ₃ , HDO
3753.0	86.0	H ₂ O, CO ₂ , O ₃ , N ₂ O, HF, HDO
3798.8	7.46	H ₂ O, CO ₂ , CH ₄ , O ₃
3822.5	51.0	CO ₂ , H ₂ O, N ₂ O, O ₃ , CH ₄ , HF, HDO
3869.14	42.0	H ₂ O, CO ₂ , O ₃ , N ₂ O, CH ₄ , HF, HDO
3895.36	10.18	H_2O , CO_2 , CH_4 , O_3
3903.93	6.17	H_2O , CO_2 , CH_4 , O_3
3936.15	58.0	CO ₂ , H ₂ O, N ₂ O, O ₃ , CH ₄ , HF, HDO
3967.79	2.0	H_2O , CO_2 , CH_4 , O_3
4026.0	114.0	H ₂ O, CO ₂ , O ₃ , N ₂ O, CH ₄ , HF, CO
4085.69	3.4	HDO, CO, H ₂ O, O ₃ , CH ₄
4115.21	5.62	HDO, CO, O_3 , CH_4
4128.34	9.53	HDO, CO, H ₂ O, O ₃ , CH ₄
4132.1	72.5	CO, H ₂ O, CO ₂ , CH ₄ , O ₃
4214.2	94.1	CO, H ₂ O, CO ₂ , CH ₄ , O ₃
4300.4	76.6	CO, H ₂ O, CH ₄ , O ₃
4377.0	75.4	H ₂ O, O ₃ , CH ₄ , N ₂ O, CO
4436.2	42.6	CO ₂ , H ₂ O, CH ₄

be seen increasing towards higher wavenumbers in Fig. 2, even after averaging the ACE-FTS spectra.

The secondary panels in Fig. 2 show the mean residuals for each fitting window at each pressure level. Shown are mean residuals when fitting with HITRAN 2012, GEISA 2015, and HITRAN 2016. HITRAN 2012 tends to have the largest numbers of errors and both newer line lists improve upon it. The most significant errors are in the CO_2 band centred at 3508 cm⁻¹, which were corrected in the 2016 edition (and not present in GEISA 2015). Difficulty in fitting strong CO_2 lines in the region between 3500–3600 cm⁻¹ persists when using all three line lists, but this is also due to a strong increase in detector noise in that region, visible in the mean spectra, especially at the lowest pressure level.

In general, HITRAN 2016 performs better than GEISA 2015, and larger mean residuals are seen in Fig. 2. The largest differences are observed the lowest pressure levels, while both line lists perform similarly at higher tangent altitudes. This suggests that the errors are related to either increased pressure and pressure broadening at lower altitudes, or (and) that the magnitude of the errors are related to line depth, and that when the line depths are small, the errors are less than the instrument noise. When CH₄ or CO₂ lines are saturated, the residuals show the negative impact of untreated line mixing, which was not implemented in this spectral region for this study, but is available in GGG [45]. Key areas that GEISA 2015 has difficulty with relative to HITRAN 2016 are between 2700–3200 cm^{-1} , when strong CH₄ lines absorb totally, between 4100-4300 cm^{-1} , where errors are related to O_3 lines and also CH₄ lines (shown in the following section). These are also errors near 3800 cm⁻¹ related to H₂O. Fitting the strong CO₂ lines between 3500-3700 cm⁻¹ is challenging for GGG using any line list, as shown in Fig. 2, especially the middle pressure level. The ν_1 band of HNO₃ is located in this region which is significant at 20 km.

At lower wavenumbers, where detector noise is lowest, there are systematic features observable in the mean residuals when



Fig. 1. A look at ACE-FTS observations in the ExoMars spectral region of interest: (a) mean ACE-FTS spectra at the 0.052 atm pressure level (near 20 km), the colours indicate different fitting windows; (b) contributions from CO_2 ; (c) contributions from H_2O ; (d) contributions from CH_4 ; (e) contributions from O_3 and N_2O .

using either line list. Near 2550 cm⁻¹, this is related to line widths in an N₂O absorption band. Near 2650 cm⁻¹, there is an observed absorption band of HNO₃ that is not included in either line list. Several HNO₃ bands are observable in ACE-FTS spectra that are not contained in the HITRAN or GEISA line lists. Weaker bands result in an apparent baseline curvature, while stronger bands show a distinct vibration-rotation band structure in the residuals. Examples are shown in the following section. Fig. 3 shows the mean RMS for each fitting window, at each pressure level. The RMS was computed for each ACE-FTS occultation in a given window, and shown are the means of the RMS values, with the standard deviation of the mean. At the highest pressure level (top panel, near 60 km), all three line lists perform similarly. In the middle pressure level, near 40 km, we observe strong increases in the mean RMS values, and their uncertainties, where line depths extend beyond 50%, just above 3500 cm⁻¹ due



Wavenumber (cm⁻¹)

Fig. 2. Mean ACE-FTS spectra, mean best-fit computed spectra, and mean residuals for each fitting window. Mean residuals are shown using HITRAN 2012 (green), HITRAN 2016 (blue) and GEISA 2015 (orange). Mean fitted spectra are shown only for when using HITRAN 2016. (a) mean residuals from 0.00023 atm (\sim 60 km), (b) mean spectra and mean fits from 0.00023 atm, (c) mean residuals from 0.0029 atm (\sim 40 km), (d) mean spectra and mean fits from 0.0029 atm, (e) mean residuals from 0.052 atm (\sim 20 km), (f) mean spectra and mean fits from 0.052 atm. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

to CO_2 , and near 3000 cm⁻¹ due to CH_4 . At the lowest pressure level, the RMS uncertainties increase strongly for all windows, and the differences between line lists become most apparent. At this level, we see deviations between HITRAN 2012 and 2016, at the same locations observed in Fig. 2. This analysis also supports the observations made from Fig. 2 regarding GEISA 2015. Specific examples and windows are explored in the following section.

6. Specific examples

6.1. 2440 – 2660 cm⁻¹: N_2O and HNO_3

The spectral region between 2440 – 2660 cm⁻¹ was covered by four broad fitting windows centred at 2455.5, 2491.5, 2551.55, and 2615.74 cm⁻¹. This region is characterized primarily by the *P* and



Fig. 3. Mean RMS values computes for each fitting window when using each HITRAN 2012, HITRAN 2016, and GEISA 2015 at three pressure levels: (a) 0.00023 atm (\sim 60 km), (b) 0.0029 atm (\sim 40 km), (c) 0.052 atm (\sim 20 km).

R branches of two vibration-rotation bands of N₂O, centred near 2462 and 2562 cm⁻¹, and two vibration-rotation bands of CO₂, centred near 2501 and 2614 cm⁻¹. For these bands, all three spectroscopic line lists perform similarly. In the 2491.5 cm⁻¹ window, small, but visible, improvements are apparent where CO₂ lines are present in the two updated editions of the line lists when compared to HITRAN 2012.

Both the GEISA 2015 and HITRAN 2016 line lists perform almost equivalently in this region. However, there are apparent problems in the fitting, and these problems persist in both data sets. At the lower pressure level, the N₂O fits produce large mean residuals whose shape is indicative of errors in line width. These errors do not persist at the higher pressure levels and are likely due to errors in the broadening parameters in the line lists. The mean ACE-FTS spectra and mean residuals for the 2551.55 cm⁻¹ window, which are shown in Fig. 4. The lines in this window are mainly due to N₂O, with some CH₄ lines throughout, and the edge of an *R* branch of CO₂ on the left side.

Fig. 5 shows the same information as Fig. 4, but for the 2615.74 cm⁻¹ window. Again, both the GEISA 2015 and HITRAN 2016 line lists perform equivalently, and, again, there are systematic differences between the observations and the computed spectra. These differences are due to HNO₃ lines not included in ei-

ther line list. The broad curvature of the baseline in the residuals is due to a weaker band, while the distinct peaks to the right side of Fig. 4 are the *P* branch of a stronger HNO₃ vibration-rotation band. There is a corresponding *R* branch in the adjacent 2666.6 cm⁻¹ window. The feature of another weak HNO₃ band is also present in the 2491.5 cm⁻¹ window, and stronger missing lines are seen in the 2978.2 cm⁻¹ window, between 2985–3010 cm⁻¹, and in the 3391.15 cm⁻¹ window, between 3390–3415 cm⁻¹. That these are due to HNO₃ is verified by measurements made for the line list distributed by the Pacific Northwest National Laboratory (PNNL) [31,55]. Supplementary HNO₃ line lists are used for ACE-FTS and TCCON retrievals.

6.2. 2660 – 3440 cm^{-1} : CH_4

This wavenumber region is characterized by strong CH_4 absorption features, but also contains important bands of O_3 , N_2O and CO_2 . This region was covered by several narrow windows and 12 wide windows ranging in width between 18–108 cm⁻¹. The first window centred at 2722.25 cm⁻¹ features three strong CH_4 lines, and many associated weaker lines, and there are significant residual errors about each due to unaccounted line mixing. Toon et al. [60] pointed out a positional error in a CH_4 line at



Fig. 4. (a) Mean residuals for the 2551.55 cm⁻¹ window at 0.052 atm (\sim 20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. (b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). The primary features in this window are N₂O lines and the residuals correspond to the N₂O lines.



Fig. 5. (a) Mean residuals for the 2615.74 cm⁻¹ window at 0.052 atm (\sim 20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. (b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). The main features are from N₂O towards the left edge, a CO₂ band across the entire window, and several strong CH₄ lines. The residuals are due to HNO₃.

2742.3 cm⁻¹ that was introduced into HITRAN 2012 after the 2008 version. This error is present in GEISA 2015 and HITRAN 2016, and is shown in Fig. 6a. While positive changes have been made to the HITRAN 2016 line list, correcting the positional error, strong residual features remain, but these are most likely due to unaccounted line mixing for such strong lines fit at a low altitude. This window also features several weak O₃ lines that are well fit by both line lists, except a small region around 2705–2710 cm⁻¹ where fitting with GEISA 2015 results in larger mean residuals than fitting with HITRAN 2016, on the order of 0.005.

The window centred at 2780.74 cm⁻¹ contains weaker CH_4 lines and a large O_3 band. Both line lists perform similarly here, but there is what appears to be a single missing line in the GEISA

2015 list. This is an ozone line at 2773.15 cm⁻¹ that is not missing, but has an error in line strength, shown in Fig. 6b. Toon et al. [60] pointed out an error in position in an O₃ resonance transition at 2761.42 cm⁻¹ that is seen in HITRAN 2012 and GEISA 2015, but has been corrected in HITRAN 2016. However, there are new O₃ positional errors at 2763.86 and 2798.0 cm⁻¹ in the HITRAN 2016 line list that were not present in HITRAN 2012.

The fitting window centred at 2825.0 cm⁻¹ presents one of the largest discrepancies between HITRAN 2016 and GEISA 2015. The mean ACE-FTS spectra and mean residuals when fitting with the three line lists are shown in Fig. 7 for the lowest pressure level. These discrepancies are characteristic of the strong CH₄ lines throughout this region, and the systematic differentiation between



Fig. 6. Zoom of the residuals and mean spectra (as in Figs. 4 and 5) where errors are present, for the windows: (a) 2722.25 cm⁻¹, (b) 2780.74 cm⁻¹, (c) 3869.14 cm⁻¹, (d) 3936.15 cm⁻¹.



Fig. 7. a) Mean residuals for the 2825.0 cm⁻¹ window at 0.052 atm (\sim 20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). Features are primarily CH₄ lines.

fitting results using the two line lists is seen in Fig. 3. There is little difference between the two versions of HITRAN, and both result in mean residual errors on the order of ± 0.01 about strong CH₄ lines. In several positions, however, when fitting with GEISA 2015, these residuals can be twice as large.

Windows centred at 2849.15, 2904.43, 2978.2, and 3126.65 cm⁻¹ are similar to those preceding, predominantly featuring strong CH₄ lines, with line mixing errors apparent in the mean residuals when using either line list, but the largest residuals are found when using GEISA 2015. In the 2978.2 cm⁻¹ window, there are several saturated (at low altitude) CH₄ lines, and the beginning of a broad, strong band of O₃ lines, but also a set of HNO₃ lines that are missing in both line lists.

The 3022.13 cm⁻¹ window contains a mixture of strong O_3 and CH₄ lines. There is an opaque region about 5 cm⁻¹ wide at the lower pressure level, which is the Q-branch of this CH₄ band. This feature is critically important for ExoMars, as it is the strongest and broadest CH₄ feature in the available wavenumber range, and where the ExoMars instruments will focus their search for a CH₄ signature. Mean residuals on the right side of the Q-branch (3018 cm⁻¹) are nearly equivalent when using GEISA 2015 or HI-TRAN 2016. On the left side (3015 cm⁻¹), mean residuals when using GEISA 2015 are larger than when using HITRAN 2016 by 0.01–0.02. Most of the fitting differences in this region can be attributed to CH₄, but there is also a weaker HCl line at 3014.4 cm⁻¹ that contributes to these differences. Note that the residuals for both



Fig. 8. a) Mean residuals for the 3089.75 cm⁻¹ window at 0.052 atm (\sim 20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). Absorption are due to, in order of prominence, CH₄, H₂O, and O₃. Large residuals near 3090 cm⁻¹ are from CH₄ lines.

line lists are significant (~ 0.002) even at the highest pressure level (near 60 km), where line depths are much weaker and only extend to 0.9. Of particular interest to ExoMars is the recently observed, weak CO₂ band centred near 3000 cm⁻¹ that overlaps with these strong CH₄ features [9,67]. This band is not visible in terrestrial solar occultation observations, its lines are too weak relative to the abundant absorption features of CH₄ and O₃.

Fig. 8 shows the mean spectra and mean residuals for fitting window 3089.75 cm⁻¹. This region contains a set of three CH₄ lines near 3090 cm⁻¹ that are poorly fit with either line list. When using HITRAN 2012 and HITRAN 2016, the magnitude of the mean residuals is 0.04, which is on the order of the line mixing errors seen when fitting stronger lines. When using GEISA 2015, the mean residuals are 50% larger. Fig. 8 also shows a pair of saturated CH₄ lines, and fitting errors characteristic of such lines.

The window centred at 3126.65 cm⁻¹ has some of the strongest CH₄ lines, and, therefore, the largest residuals. GEISA 2015 and HITRAN 2016 perform similarly here. The window centred at 3202.0 cm⁻¹ contains the right edge of CH₄ ν_2 band. At higher wavenumbers, H₂O absorption becomes dominant. This window also features overlapping CO₂ and O₃ bands, the edge of the strong CH₄ band, and several water vapour lines. The 3292.0 cm⁻¹ window contains H₂O and CO₂ lines, with an H₂O line strength error at 3254.15 cm⁻¹ that has been reduced between HITRAN 2012 and HITRAN 2016, but is still significant in both HITRAN 2016 and GEISA 2015. This window contains several other H₂O lines where significant improvements were made to both line lists since HI-TRAN 2012 (e.g., at 3273.4 and 3276.5 cm⁻¹). The 3391.15 cm⁻¹ window is similar, but hosts another set of HNO₃ lines missing from both line lists. There is also an H₂O line strength error at 3367.65 cm⁻¹ in both GEISA 2015 and HITRAN 2016 that was not significant in HITRAN 2012.

6.3. $3440 - 3770 \text{ cm}^{-1}$: CO₂

This region was covered by five wide windows centred at 3391.15, 3489.0, 3577.0, 3665.0, and 3753.0 cm⁻¹. These windows cover the strongest CO_2 vibration-rotation bands observed by the ExoMars instruments. Throughout this region, there are only small

differences between fits using GEISA 2015 and HITRAN 2016. However, there are significant improvements to several lines throughout since HITRAN 2012. Fig. 9 shows the mean spectra and residuals for the window centred at 3489.0 cm⁻¹, where the most dramatic improvement is seen. The spectroscopic parameters for an entire band have been updated, resulting in significant improvements to fitting.

6.4. 3770 – 4080 cm⁻¹: H_2O

This spectral region, covered by four broad fitting windows centred at 3822.5, 3869.14, 3936.15, and 4026.0 cm⁻¹, features the strongest set of water vapour lines in the ExoMars region of interest. It is in this region where we again find that spectral fitting with the HITRAN 2016 results in smaller residuals than when using GEISA 2015, as seen in Fig. 3. Fig. 10 shows the window centred at 3869.14 cm⁻¹. The primary features are broad, saturated H₂O lines, but there are also many smaller CH₄ and N₂O lines, and many very weak lines from CO₂ and O₃. The lines where fitting with GEISA 2015 resulted in systematic residual errors are largely attributable to weaker O₃ lines. Similar behaviour is seen in the window centred at 3822.5 cm⁻¹, where several O₃ lines are not contained in GEISA 2015. Note the marked improvement in the strong H₂O line between HITRAN 2012 and the newer line lists near 3886 cm⁻¹, as shown in Fig. 6c, but clearly visible in Fig. 10.

The two windows at 3936.15 and 4026.0 cm⁻¹ show improvements over the previous two windows when comparing GEISA 2015 and HITRAN 2016. The 3936.15 cm⁻¹ window contains a portion of an O_3 *P*-branch. Significant residuals when using GEISA 2015 are from widely spaced CH₄ lines (e.g., at 3912.15, 3914.6, and 3914.9 cm⁻¹), rather than O_3 . Two of these features near 3914.6 cm⁻¹ are shown in Fig. 6d. There is an H₂O line strength error in the HITRAN 2016 line list at 3925.15 cm⁻¹ that was not present in HITRAN 2012, nor GEISA 2015. The 4026.0 cm⁻¹ window hosts an O_3 band, for which all line lists have difficulty accurately fitting over the region of the Q-branch. The O_3 band in this region is much better characterized by GEISA 2015 than that covered by the 3822.5 and 3869.14 cm⁻¹ windows, but there remain



Fig. 9. (a) Mean residuals for the 3489.0 cm⁻¹ window at 0.052 atm (\sim 20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. (b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). The primary features are an N₂O band on the left side, CO₂ bands on the right, and several H₂O lines. The strong residuals are due to CO₂ when using the HITRAN 2012 line list.



Fig. 10. a) Mean residuals for the 3869.14 cm⁻¹ window at 0.052 atm (~20 km) with HITRAN 2012, HITRAN 2016, and GEISA 2015. b) Mean measured ACE-FTS spectrum and mean computed spectrum (for HITRAN 2016). The primary features are H_2O lines. The strong residuals near 3860 cm⁻¹ are due to O_3 when using GEISA 2015. Those near 3886 cm⁻¹ are due to H_2O and are shown in Fig. 6c.

small discrepancies in some minor CH_4 lines, especially towards the right edge of the window at 4080.65 and 4082.8 cm⁻¹.

6.5. 4080 - 4420 cm⁻¹: CH₄

Above 4100 cm⁻¹, the ACE-FTS noise increases, as can be seen in Fig. 2, and it becomes difficult to evaluate the line lists further. This region is characterized by a decreasing density of H_2O lines, increasingly stronger, and more dense CH_4 absorption features, and a band of broadly spaced CO lines. We evaluated five wide fitting windows in this region centred at 4132.1, 4214.2, 4300.4, 4377.0, and 4436.2 cm⁻¹. At 4115.65 cm⁻¹, a CH_4 line missing in HITRAN 2012, as identified by Toon et al. [60], was added (also included in GEISA 2015). There are CH_4 line strength errors in the GEISA 2015 line list at 4133.35 cm⁻¹ and 4103.2 cm⁻¹ that result in mean residuals on the order of 0.1 at lower pressure levels. The latter is also apparent in HITRAN 2016, despite not being present in HI-TRAN 2012.

As shown in Fig. 3, the mean RMS values dramatically increase in this spectral region for all evaluated line lists. At the lowest pressure level, fitting with HITRAN 2016 results in the smallest residuals and the lowest mean RMS values. It must be noted that the magnitude of the difference between results using HITRAN 2016 and GEISA 2015 is on the order of the difference between changing pressure levels, and that is only \sim 1/5 of the magnitude of the mean RMS. Therefore, errors due to differences in the spectroscopic parameters are much, much smaller than the noise level of the instrument in this region.

The window centred at 4214.2 cm⁻¹ features the most dramatic difference between fitting using GEISA 2015 and fitting using HI-TRAN 2016. Mean residuals with magnitudes greater than 0.1 come from several CH₄ lines near 4208, 4229, and 4255 cm⁻¹.

7. Conclusions

This study was motivated by the release of two new editions of spectroscopic line lists, the 2015 version of GEISA and the 2016 version of HITRAN, and the launch and arrival of the ExoMars Trace Gas Orbiter as Mars, equipped with two suits of spectroscopic instruments dedicated to characterizing the Martian atmosphere. The largest efforts made recently towards updating infrared spectroscopic databases has been in support of terrestrial greenhouse gas observatories such as OCO-2, GOSAT, and TCCON. Since the Martian atmosphere in composed of 96.5% CO₂, these updates are very significant for the ExoMars mission. Our objective was to validate the two line lists in the range of 2325–4350 cm⁻¹ by examining spectral fitting results for terrestrial solar occultation observations made by ACE-FTS.

This work follows that of [60] who compared different versions of HITRAN, up to the 2012 release, in the spectral range of 670– 5620 cm⁻¹ using the solar occultation measurements made by the MkIV FTS. They identified several errors in HITRAN 2012, some persisting from previous releases, and others newly introduced. The analysis in [60] is expanded in [58], which includes a detailed analysis of HITRAN 2016 using laboratory and solar occultation spectra. Because spectroscopic parameters taken from previous HITRAN versions for some gases in some spectral regions perform better, and because HITRAN may be incomplete in some spectral regions (*e.g.*, HNO₃), the TCCON and GGG development teams maintain a custom line list, as do other spectroscopic analysis teams, such as ACE-FTS.

The 2016 edition of the HITRAN line list addressed several errors identified in the 2012 edition by Toon et al. [60], such as the positional error in an O_3 resonance transition at 2761.42 cm⁻¹ (still persists in GEISA 2015). We have observed, however, a few minor errors introduced into HITRAN 2016 since the 2012 edition, such as the H₂O line strength errors at 3367.65 cm⁻¹ (also in GEISA 2015) and 3925.15 cm⁻¹.

For ExoMars, the gases of primary interest are CO_2 , H_2O , CO, and CH_4 . Changes to line position and strength have been made to CO_2 lines across our spectral region. For terrestrial spectra, the strongest improvements seen in fitting coincide with the strongest absorption features, and the greatest improvement is seen between 3470–3530 cm⁻¹ for both GEISA 2015 and HITRAN 2016. In no region were increased residuals seen for CO_2 lines when using the updated line list editions when compared to HITRAN 2012. In the terrestrial observations, we also observe significant improvements to a subset of H_2O lines, especially at higher wavenumbers, such as near 3093.7 or 3885.5 cm⁻¹. With the exception of individual CH_4 line errors, little difference is observed in the CH_4 or CO transitions in this spectral region between HITRAN 2012 and 2016.

Of key interest to us was the CO_2 vibration-rotation band centred at 2982 cm⁻¹ and partially overlapping the critically important CH₄ ν_2 band. When comparing synthetic spectra generated with HITRAN 2016 or GEISA 2015, significant differences in these lines are seen. This band is absent from GEISA 2015, but present in HITRAN 2012 and 2016. The lines included in HITRAN 2016 have significantly increased line strengths relative to HITRAN 2012. Unfortunately, in terrestrial observations, these lines are too weak relative to interfering species, especially CH₄, and noise. It is not ob-

served in the ACE-FTS spectra examined here above background noise levels, so residuals between spectral fits using HITRAN 2012 and 2016 have not been compared.

A critical difference between the application of spectroscopic calculations for Earth and Mars is that the because the Martian atmosphere is predominantly CO_2 rather than N_2 , the collision-induced broadening parameters computed for HITRAN and GEISA will not be applicable to the Martian atmosphere. There is ongoing work to determine spectroscopic parameters for a CO_2 -rich atmosphere: *e.g.*, [20,24] for water vapour; and [38] for CO. However, this study does not attempt to validate these parameters, nor does it evaluate the CO_2 self-broadening parameters.

When comparing the 2015 version of the GEISA line list to the latest, 2016, release of HITRAN, we observe that lower RMS values are found for residuals from the majority of the spectral windows between 2325–4350 cm⁻¹ used here. We find that these are primarily due to differences in line strength or position for strong O_3 and CH₄ lines. There are some minor errors in specific lines noted as well.

We were surprised to find a large number of O_3 lines missing from GEISA 2015 in the 3830–3870 cm⁻¹ region that were present in HITRAN 2012. The primary source of O_3 lines both HITRAN 2016 and GEISA 2015 is the Spectroscopy and Molecular Properties of Ozone (S&MPO) information system [2] maintained by Reims University and the Institute of Atmospheric Optics (Tomsk). For GEISA 2015, new measurements of lines around this region were made at Reims University by Barbe et al. [4,6], but those lines precisely between 3830–3870 cm⁻¹ are attributed to a private communication from Barbe (2011). O_3 line parameters in HITRAN 2016 are also included in the S&MPO database. In the region 3623–4229 cm⁻¹, HITRAN also includes an updated hot band from Barbe et al. [5].

Our analysis fit the hydrogen halides HCl (between 2600– 3050 cm⁻¹) and HF (between 3700–4100 cm⁻¹). Parameters for these gases were not updated in the 2011 or 2015 editions of GEISA, but new calculations were implemented in the 2012 edition of HITRAN [37]. The result is differences in line position when performing spectral fitting with HITRAN 2016 or GEISA 2015. Mean differences are very small relative to the noise of the ACE-FTS observations, less than 1%, with HITRAN 2016 performing slightly better at these line locations (*e.g.*, 2925.9 cm⁻¹ or 2944.93 cm⁻¹ for HCl, and 3877.7 cm⁻¹ or 3920.3 cm⁻¹ for HF).

In conclusion, we have noted consistent improvement in both line lists since HITRAN 2012, have noted errors and deficiencies in both line lists where found, and we hope that the collaborations in charge of both line lists will find this analysis useful when compiling the next release. However, for the purposes of the Exo-Mars TGO ACS and NOMAD instruments, we recommend HITRAN 2016 for use. HITRAN 2016 shows marked improvement over HI-TRAN 2012 in CO₂ and H₂O transitions (as does GEISA 2015). GEISA 2015 currently produces larger residuals for strong CH₄ lines that are critical for ExoMars, while an important CO₂ band centred at 2982 cm⁻¹ has not yet been introduced.

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Supplementary material

Supplementary material associated with this article can be found, in the online version, at https://doi.org/10.1016/j.jqsrt.2019. 106590.

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